Prediction of submerged landslide generated waves in dam reservoirs: an applied approach

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Abstract

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cities. amplitude in dam reservoir using the landslide geometry parameters and water body to provide an engineering order Boussinesq-type numerical model is applied to study the sensitivity of the wave importance in dam engineering and water reservoir planning. In this work, a higher enon may be endangering the dam stability, human life and devastating downstream flow, including debris flows, debris avalanches, landslides, and rock falls. This phenom-Impulsive waves in dam reservoirs may be generated by any type of geophysical mass amplitude to the landslide geometry and kinematics. with several experimental and real cases. conditions. The estimation of the landslide generated The estimation method is verified and proven by comparison of the results estimation method to predict the landslide waves amplitude is of the utmost The numerical results are also used generated wave

Keywords: Dam reservoir, submerged landslide, impulsive waves, landslide, Boussinesq model, Numerical model

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. Introduction

Underwater landslides in artificial reservoirs can trigger local impulsive waves with high run-up, endangering dam stability, human life and devastating downstream cities. Submarine landslides, which often accompany large earthquakes, can disturb the overlying water column as sediment and rock slump down slope. Any sort of geophysical mass flow, including debris flows, debris avalanches, landslides, and rock falls can generate submarine landslide impulsive waves. Estimating the amplitude of Landslide (ienerated Waves (LGW) is of the utmost importance in dam engineering and water reservoir planning. There are a several real cases in which the LGWs destroy the dam or the villages around it. For instance a landslide on October 9, 1963 with a volume of 270 million m¹ generated an impulse wave initially more than 70m tall in Vaiont dam reservoir in Italy, the worlds' tallest double thin arch dam. The impulse wave completely

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destroyed five villages and about 2500 people were killed. impulse waves generated by landslides are known from Lituya Bays, Alaska [18]; Yanahuin Lake, Peru; and Shimabara Bay in Japan [21]. Other examples of large

predict the impulse wave characteristics. An overview of numerical modelling of landslide waves is presented in Table 1. A major part of this body of works is related to instability. As it can be seen, the submarine landslide generated tsunami is fairly well landslide tsunami waves which are a special type of tsunami waves caused by slope investigated but the application of impulse wave models for forecasting the submerged numerical studies, experimental investigations have been performed with regards LGW characteristics in dam reservoirs still requires further studies. In addition to the generated tsunami. The basic parameters can be seen in this table that most experimental works are related to submarine landslide LGWs. An overview of laboratory works and initial submergence of failure mass merged body are set in accordance with landslide tsunamis. Moreover, the conclusion of empirical formulations is specially made and verified for tsunami applications. For instance, the recent studies on landslide generated tsunamis performed by Grilli et al. in ed by Panizzo et al. in 2005 [22] the maximum bed slope is fixed on $\theta=36^\circ$ specification of landslide and characteristics of water body when using landslide tsunafocused on sub-aerial landslide. There 2005~[8] has a main constraint for bed slope as $\theta < 30^\circ$ or in other recent works presenting the main characteristics of impulse mi predictive methods in a dam reservoir. Thus, more research is required for forecastwhere the bed slope may be much steeper than the inclined sea floor and the depth, location of failure mass, propagation distances and water body specifications are seriously different from ocean. The numerical modelling and experimental investigations have are noticeable constraints in other geometrical in experiments consist of bed slope, geometry related to this topic is presented in Table 2. It and particularly the laws of motion of wave caused by landslide in dam reservoir, been developed to

simple engineering estimation method for prediction of impulse wave amplitude caused by underwater landslide in dam reservoirs. The main information required consists of landslide geometry and water body conditions. The sensitivity analysis of wave amplimodel is an extension of (4,4) Padé approximant [4] to include moving bottom boundary to be applicable to study LGWs. The details of derivation for extended formulations sented in Reference [1]. In this work, an extended higher order Boussinesq-type model is used to provide well as also made here using numerical numerical modelling, calibration and verification of model are clearly details of derivation for extended formulations The Mathematical formulation of the

Mathematical formulation

based on fourth order Boussinsq-type e The mathematical formulation of the numerical model used in this work is developed quations, known as (4,4) Padé approximant

No.	Reference	Ref	Numerical	Governing	Numerical	Model dimensions	Failure Mass	Wave simulation	Experimental	Real case study
			Model Name	equations	method			stage	validation	
1	Noda (1971)	[19]	_	NSW	FDM	Depth averaged	Rigid sliding	Generation	Done	
2	Mader (1973)	[17]	ZUNI	NSW	FDM	Width averaged	Rigid sliding	Gen., Prop., run-up	Done	
3	Raney/Butler (1975)	[25]	(<u>5</u>	NSW	FDM	Depth averaged	Rigid sliding	Gen., Prop., run-up	_	Landslide Tsunami
4	Goto/Ogawa (1992)	[5]	TIME	BW	FDM	Depth averaged	Rigid sliding	Gen., Prop., run-up		Landslide Tsunami
5	Jiang/Leblond (1992)	[12]		NSW Viscous flow	FDM	Three dimensional	Mudslide	Gen., Prop.	Done	Landslide Tsunami
6	Titov (1997)	[29]	MOST	NSW	FDM	Depth averaged	Rigid sliding	Gen., Prop., run-up	£	Landslide Tsunami
7	Rzadkiewicz et al	[26]	Nasa-Vof2d	Potential flow	FDM	Width averaged	Slumping	Generation	Done	-
	(1997)			Bingham					± :	
			€	Rheology						
8	Grilli et al (1999)	[9]	GW-2D	Potential flow	BEM	Width averaged	Rigid sliding	Generation	Done	Landslide Tsunami
9	Imran et al (2001)	[11]		NSW	FDM	One dimensional	Slumping	Generation	Done	-
5				Herschel-						
				Bulkley						
		2 8		Rheology						
10	Synolakis et al (2002)	[28]	TUNAMI	LSW	BEM	Depth averaged	Rigid sliding	Gen., Prop.		Landslide Tsunami
11	Grilli et al (2002)	[6]	BIEM	Potential flow	BEM	Three dimensional	Rigid sliding	Gen., Prop.	Done	
12	Watts et al (2003)	[32]	GEOWAVE	BW	BEM	Three dimensional	Sliding/slump.	Gen., Prop., run-up	Done	Landslide Tsunami
13	Grilli/Watts (2003)	[7]	BING	BW	BEM	Width averaged	Slumping	Gen., Prop.	Done	<u></u>
				Bingham Rhe.						
14	Lynett/Liu (2004)	[16]	COULWAVE	BW-Multi	FDM	Depth averaged	Slumping	Gen., Prop., run-up	Done	Landslide Tsunami
3				Layer						
15	Panizzo et al (2005)	[22]		Potential flow	SPH	Three dimensional	Slumping	Gen., Prop.	Done	Landslide wave in res.
16	Ataie/Jilani (2006)	[1]	953 95 (655/869/668) 16. 1 <u>4</u> 26	BW	FDM	Depth averaged	Sliding/slump.	Gen., Prop.	Done	

caused by landslide

Gen: Generation

BEM: Boundary Element Method FDM: Finite Difference Method SPH: Smooth Particle Hydrodynamics

NSW: Nonlinear Shallow Water Wave Equations

LSW: Linear Shallow Water Wave Equations

Prop: Propagation

BW: Boussinesq-type Wave Equations

 $w = \mu^2 w_1 + \mu^4 w_2$

 \otimes

 $= u_0 + \mu^2 u_1 + \mu^4 u_2$

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step of the Lynett & Liu approach [15] which was used for extension of second-order Boussinsq wave equations. A schematic of the main geometric The model includes the moving bottom boundary. The extension procedure is a forward three-dimensional domain can be described as [15]: The dimensionless form of governing equations and boundary conditions in parameters is shown in

$$\mu^{2}\nabla.u + w_{z} = 0 \qquad \text{on } -h \le z \le \varepsilon. \zeta \quad \text{(Continuity equation)}$$

$$u_{t} + \varepsilon u \nabla.u + \frac{\varepsilon}{\mu^{2}} w u_{z} = -\nabla p \qquad (1 + \varepsilon u \nabla.u + \frac{\varepsilon}{\mu^{2}} w u_{z} = -\nabla p \quad (2 + \varepsilon.) \zeta \quad (3 + \varepsilon.)$$

on

-h ≤ 2 ≤

(Continuity equation)

on
$$-h \le z \le \varepsilon . \zeta$$
 (Momentum Equation in 2 horizontal dir.)

$$\varepsilon w_t + \varepsilon^2 u.\nabla w + \frac{\varepsilon^2}{\mu^2} w w_z = -\varepsilon P_z - 1$$
on $-h \le z \le \varepsilon.\zeta$

(Momentum equation in z dir.)

(4)

(5)

on z =8.5 (KFSBC)

 $w = \mu^2 \left(\zeta_t + \varepsilon. u. \nabla \zeta \right)$

(DFSBC)

 $v + \mu^2 u \cdot \nabla h + \frac{\mu^2}{\varepsilon} h_t = 0$

length scale, z is the vertical coordinate scaled by h_0 which is the characteristic

is the water surface displacement scaled by

where x and y are the horizontal coordinates scaled by t_0 which is the horizontal

 a_0 which is the impulse wave amplitude, h is the total depth based on still water consid-

depth, t is time and scaled by $l_0/(gh_0)^{1/2}$

ering the moving bottom boundary (h(x,y,t))

zontal velocity components (u,v) scaled by $(\varepsilon(gh_0)^{1/2}, w)$ is the velocity in value (u,v) scaled by $(\varepsilon(gh_0)^{1/2}, w)$ is the velocity in value (u,v) scaled by (v,v) and (v,v) scaled by (v,v) and (v,v) is the water pressure scaled by (v,v) and (v,v) and (v,v) is the water pressure scaled by (v,v) and (v,v) and (v,v) is the water pressure scaled by (v,v) and (v,v) and (v,v) is the water pressure scaled by (v,v) and (v,v) and (v,v) is the water pressure scaled by (v,v) and (v,v) and (v,v) is the water pressure scaled by (v,v) and (v,v) and (v,v) is the water pressure scaled by (v,v) and (v,v) and (v,v) is the water pressure scaled by (v,v) and (v,v) and (v,v) is the water pressure scaled by (v,v) and (v,v) and (v,v) is the water pressure scaled by (v,v) and (v,v) and (v,v) is the water pressure scaled by (v,v) and (v,v) and (v,v) is the water pressure scaled by (v,v) and (v,v) and (v,v) is the water pressure scaled by (v,v) and (v,v) and (v,v) is the water pressure scaled by (v,v) and (v,v) and (v,v) is the water pressure scaled by (v,v) and (v,v) and (v,v) is the water pressure scaled by (v,v) and (v,v) and (v,v) is the water pressure scaled by (v,v) and (v,v) and (v,v) is the water pressure scaled by (v,v) and (v,v) is the water pressure scaled by (v,v) in (v,v) is the water pressure scaled by (v,v) in (v,v) is the water pressure scaled by (v,v) in (v,v) is the water pressure scaled by (v,v) in (v,v) in (v,v) is the water pressure scaled by (v,v) in (v,v) in (v,v) in (v,v) is the water pressure scaled by (v,v) in (v,v) in (v,v) is the water pressure scaled by (v,v) in (v,v) in (v,v) is the water pressure scaled by (v,v) in (v,v)

by $(\varepsilon(gh_0)^{1/2}, w)$ is the velocity in vertical direc-

and scaled by h_0 , **u** is the vector of hori-

flow condition is assigned for lateral boundaries and the numerical simulation shall be

 a_0/h_0 and $\mu = h_0/l_0$, respectively.

partial derivative. The non-

The no-

 $-(\partial/\partial x , \partial/\partial y)$

lateral boundaries.

In perturbation

stopped before the generated waves are received by

analysis, the velocity domain components "u" and "w" shall be expanded in to [15]:

linearity and dispersion parameters are

the horizontal gradient vector. The subscripts denote the

No Reference Bed slope | Failure Mass Ref Wave Tank/Flume **Impluse** Model Wave Slide Comparison Notes No Specifications (deg) specifications Mechanism Dimensions Simulation Location with num L(m)W(m)H(m)Stage Data Jonson Bermel (1949) [13] Shallow Water Tank Steel plate Rigid Generation Sub-aerial [34] [24] Wiegel (1955) Crashing Prins (1958) Shallow Water Tank Kamphuis/Bowering [14] Rigid sliding 45 Steel box Sub-aerial Generation Definition of (1972)compact Froude No Heinrich (1992) [10]4.0 0.3 2.0 Triangle solid Rigid/ Flume/2VD Generation Submerged Done 30 block Slumping (two-vertical (50x50cm)/dimension) mass Gravel with Identical Dia Watts (1998) [31] **PVC** Triangle 9.14 0.101 0.66 45 Rigid Flume/2VD Generation Submerged Presentation of wave Solid Block Maker Curves as a (86x86cm) function of solid slide and water spec. Grilli/Watts (2005) [8] 30 3.6 1.8 30 15 Semi-Ellipse Rigid Tank /2VD Gen., Prop. Submerged Wave absorber Done 13.85 Aluminium Sliding system installed at Sheet lateral boundaries Walder et al (2003) [30] 3.0 0.285 1.0 19.5 Hollow Rect. Rigid Flume /2VD Generation Sub-aerial Done Empirical formulations 15 Nylon box Sliding for sub-aerial impulse with triangle wave estimation snout Enet et al. (2003) 30 1.8 3.6 Rigid Aluminium Tank /3D Submerged Gen., Prop. Verification of BIEM Done sheet with Sliding as a 3D potential flow Geometry of model of truncated hyperbolic secant func. 9 Watts et al (2003) [32] 30 3.6 1.8 Slumping Different Flume/2VD Generation Submerged Done granular mass materials with 3mm dia. consist of glass beads, steel shots & lead shot Fritz et al (2004) [3] 0.5 45 1.0 Failure Slumping Flume/2VD Gen., Prop. Sub-aerial Categorisation of soil mass mass impulse wave caused by caused by Pnematic landslide Landslide Generator 10 Panizzo et al (2005) [22] 11.5 0.8 Solid Rigid Tank/3D Gen., Prop. Sub-aerial Done Empirical formula. rectangular crashing and run-up for sub-aerial 36 box landslide waves in dam reservoirs

Table 2: An overview of the main laboratory investigations of impulsive waves caused by landslide

to the contribute the state of the state of

 $\begin{array}{l} h_{t} + \zeta_{t} + \nabla \cdot \left\{ (\varepsilon \zeta + h) u_{0} + \mu^{2} \left[\frac{1}{6} (\varepsilon^{3} \zeta^{3} + h^{3}) A + \frac{1}{2} \widetilde{z}^{2} (\varepsilon \zeta + h) A - \frac{1}{2} (\varepsilon^{2} \zeta^{2} - h^{2}) (\nabla \cdot B) + \widetilde{z}^{2} (\varepsilon \zeta + h) (\nabla \cdot B) \right] \\ + \mu^{4} \left[\frac{1}{120} (\varepsilon^{5} \zeta^{5} + h^{5}) \nabla (\nabla \cdot A) - \frac{1}{24} (\varepsilon \zeta + h) \widetilde{z}^{2} \nabla (\nabla \cdot A) - \frac{1}{12} (\varepsilon^{3} \zeta^{3} + h^{3}) \nabla (\nabla \cdot (\widetilde{z}^{2} \cdot A)) + \frac{1}{24} (\varepsilon^{4} \zeta^{4} - h^{4}) \nabla (\nabla \cdot (\nabla B)) - \frac{1}{6} (\varepsilon \zeta + h) \widetilde{z}^{3} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon \zeta + h) \widetilde{z} \nabla (\nabla \cdot (\nabla B)) + \frac{1}{2} (\varepsilon$

$$\begin{aligned} u_{0t} + \varepsilon (\nabla . u_0) u_0 + \varepsilon (w_1|_{z=0}) u_{0z} \\ + \mu^2 \left[u_{1t}|_{z=0} + \varepsilon (\nabla . \left(u_1|_{z=0} \right) \right) u_0 + \varepsilon (\nabla . u_0) \left(u_1|_{z=0} \right) . + \varepsilon \left(w_2|_{z=0} \right) u_{0z} + \left(w_1|_{z=0} \right) \left(u_1 z_{1} z_{2} \right) \\ + \mu^4 \left[u_{2t}|_{z=0} + \varepsilon (\nabla . \left(u_2|_{z=0} \right) \right) u_0 + \varepsilon (\nabla . \left(u_1|_{z=0} \right)) \left(u_1|_{z=0} \right) + \varepsilon (\nabla . u_0) \left(u_2|_{z=0} \right) \\ + \varepsilon \left(w_2|_{z=0} \right) \left(u_1 z_{1} z_{2} \right) + \left(w_1|_{z=0} \right) \left(u_2 z_{1} z_{2} \right) \\ + \nabla \left(P|_{z=0} \right) = 0 \left(\varepsilon^6, \mu^6 \right) \end{aligned}$$

$$(10)$$

where vector $\mathbf{A} = \nabla \cdot (\nabla \cdot \mathbf{u}_0)$, scalar $\mathbf{B} = \nabla \cdot (h\mathbf{u}_0) + h_t/\varepsilon$ and \widetilde{z} is a weighted average of two distinct characteristic water depth as described by Gobbi et al. [4]. The above equations are solved simultaneously to obtain the main three variables u, v (horizontal velocity components) and ζ (water wave elevation). A numerical model has been developed based on this set of equations using finite difference method to simulate the impulsive wave generation and propagation. The sixth order difference scheme and discretization method is applied following Gobbi et al., [4] numerical model but the moving bottom boundary is included here.

3. Wave amplitude sensitivity analysis

As shown in Figure 1, two main categories of parameters can be recognised which affected the landslide impulse waves. The first category includes parameters which are related to the geometry of the failure mass and sloping bed. These characteristics consist of bed slope; θ , the length of sliding block along the slope; B, the maximum thickness of sliding block; T, the initial still water depth at the mass center point of sliding mass; h_{0C} , and the sliding soil density; γ . The second category is the generated wave charac-

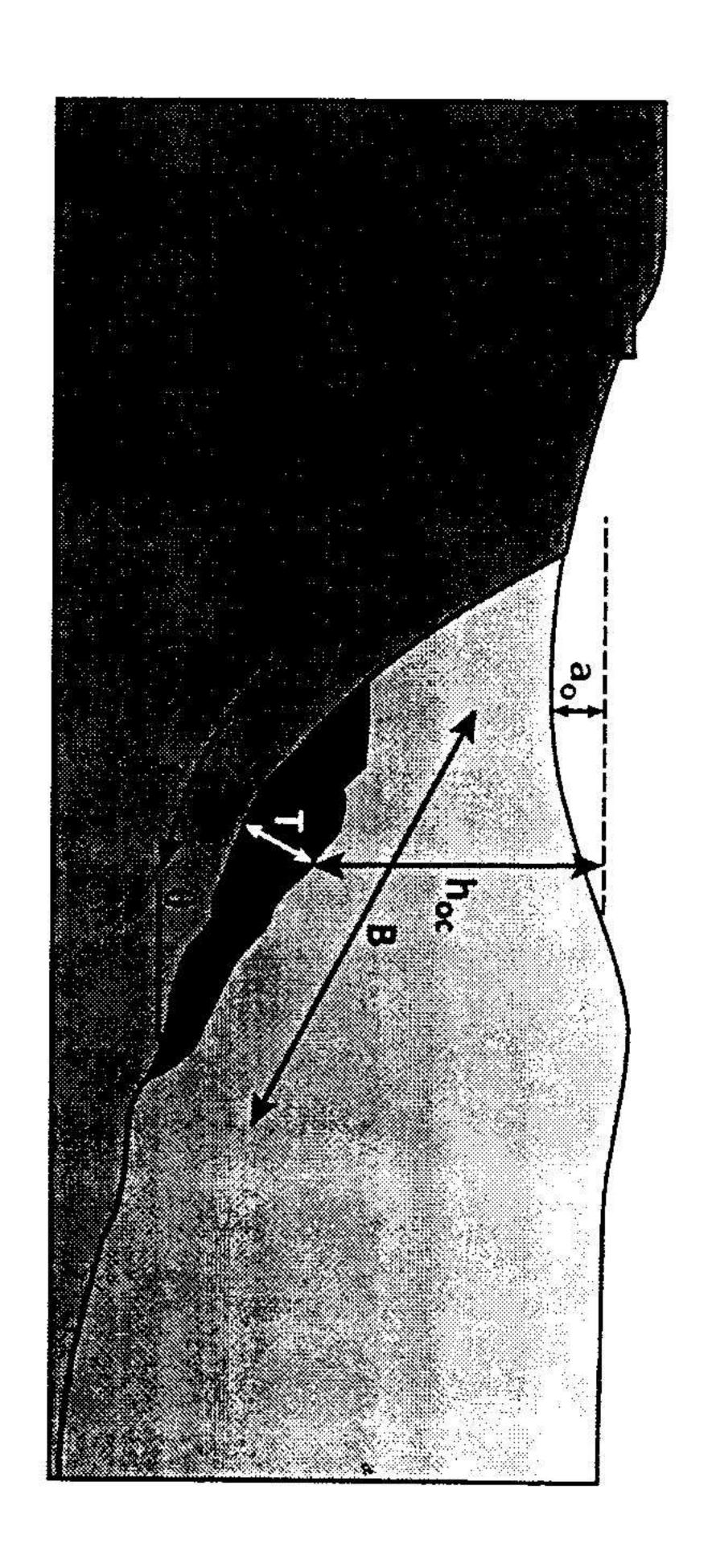


Figure 1: A schematic of main geometric parameters in submerged landslide generated impulse wave

and defined as [9]: surface (a_0) defined which describe convenient to define the basic dimensionless Maximum Another initial submergence such non-dimensional parameter can be defined u thickness (T) over the divided by a characteristic length. This characteristic length is named as S_0 Based amplitude 0n laboratory the mass (or of failure maximum length of failure investigations mass geometric depression sliding b parameters sing maximum depression of water ock along the slope (B); (i.e. T/B), over its that describe its initial location as dimensionless ratios can wavelength, length (B); (i.e. the landslide and period. follows: and

$$S_0 = \frac{u_t^2}{\alpha_0}$$
 where $u_t = \sqrt{g.B}.\sqrt{\frac{\pi(\gamma - 1)}{2C_d}}$.sin θ and $a_0 = g.\frac{\gamma - 1}{\gamma + C_m}\sin\theta$

In this equation, ut is the terminal velocity of failure mass, α_0 is its initial acceleration d d and d and d are the shape related coefficients of failure mass consist of drag and

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added mass coefficients, respectively. For meters on impulse wave amplitude, the numerical results in various conditions are presented as a set of graphs in Figure 2. In this Figure, the interaction of landslide geometry and wave amplitude can be seen. The parameters is selected based on observed underwater landslide in real cases slope; 0, consist of 5, able experimental characteristics. Some of these specifications are mentioned in section when the initial submergence of failure mass diminishes or the sliding block thickness 4. As it can be seen, in a constant value of bed slope, the wave height will be increases. The effects of the thickness of of bed slope on the impulsive wave generation and propagation are discussed. As can be and the impulsive wave amplitude will be When the bed slope increases, the sliding velocity and its acceleration will be increased seen in Figure 3, the wave amplitude increases when the bed slope enlarges. In Figure consisting of landslide geometry and kinematics as well as failure mass density are 3, the wave propagation pattern can be seen at three time stages. All of the conditions identical in all cases. The numerical wave geometric ratio of sliding block is supposed as T/B=0.1 and the initial submergence of the bed slope influences on the and for figures b, d, and f; θ =10°. In constant geometrical and kinematics specifications, sliding block ratio; $h_{0C}/B=1$. For Figures observed in Figure 3. The wave heights are multiplied by 10, for a better recognition in as wave propagation speed and wave group velocity increase in steeper slope landslide at least in near field. bed slope rather than far from it. Furthermore, it seems that the wave length as well figures. The wave height near to the source of the landslide is influenced intensively 15, 30 and 45 degrees. The range of variation of dimensionless impulsi 3a, enlarged. For further investigation, the results presented for different values of bed ve wave sliding block increases in shallower sliding. investigation of influences of various paratank dimensions assume 8x8 meters. The 3c, and 3c; the bed slope is assumed 0=5° generation and propagation can be and avail-

4. Wave amplitude prediction method

The numerical simulation results that are used for sensitivity analysis and illustrated in underwater landslide in a dam reservoir. Figure 2 can be applied for estimating the amplitude of impulsive wave caused by an mass consist of slide length along the bed slope (B), maximum thickness of failure mass to use the estimation method are motion; So which is defined in equation added mass coefficients (C_d and C_m , respectively) should be defined. Definition of the density as well as the shape the shape and dimensions of sliding mass. It must be noted that the range of validity coefficients in real case problems can be presented prediction method is γ ∈ [1.8, and initial submergence at the mass center (h_{00}). 2) The characteristic length of related coefficients of failure mass consists of drag and listed 2.1] where γ is the failure mass density here: 1) The geometric parameters of sliding (11). Based on this equation, the sliding mass The basic parameters which should be known carried out using empirical methods based on

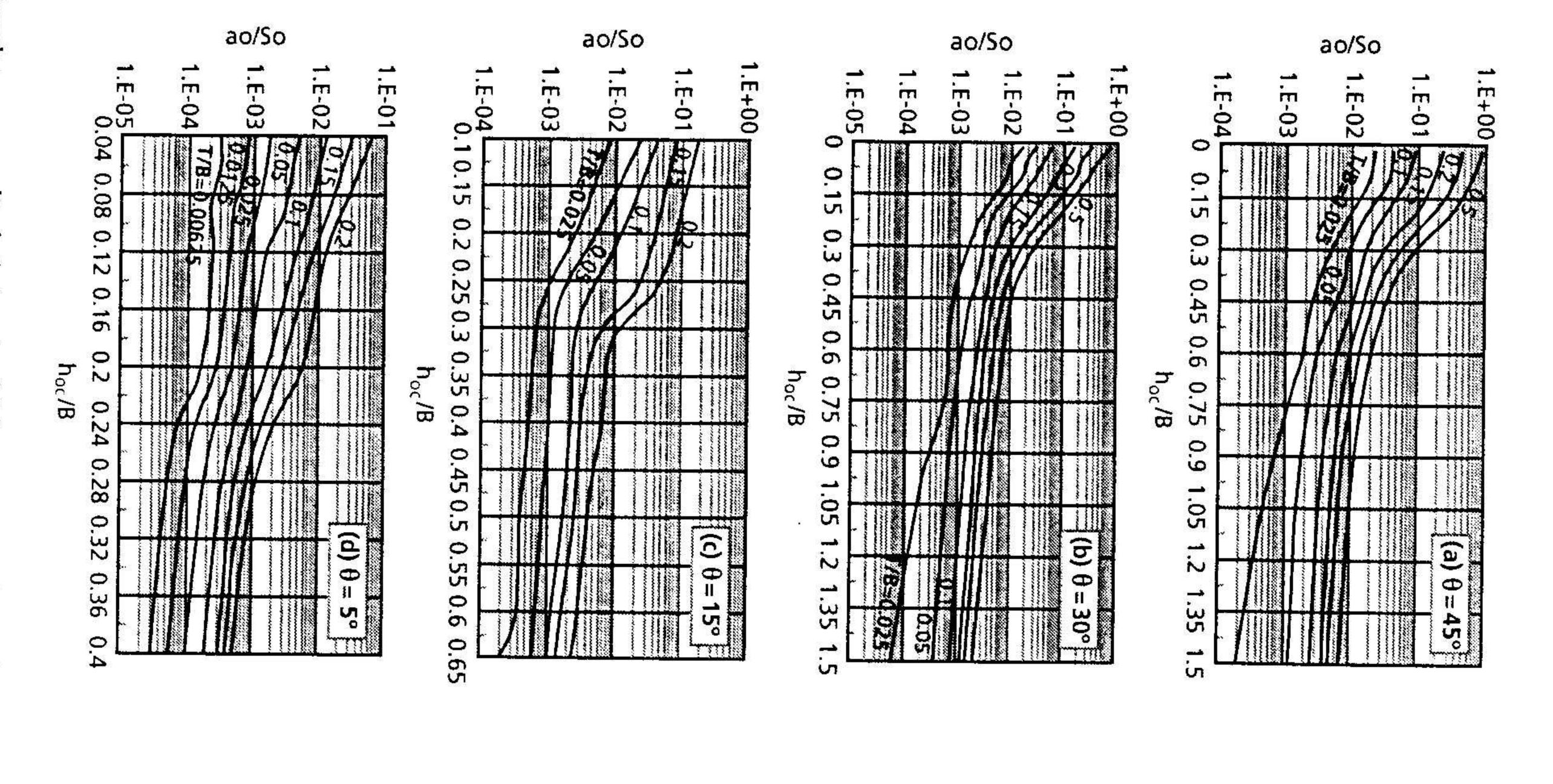


Figure 2: Impulse wave amplitude (a_o) vs. landslide geometry and water body conditions, the bed slope; θ = a)45°, b) 30°, c) 15° and d) 5°, h_{0C} is the initial submergence of landslide, B is the landslide length along the inclined bed, T is the maximum thickness of landslide and S₀ is the characteristic length of the landslidemotion which is defined in Fanation 11

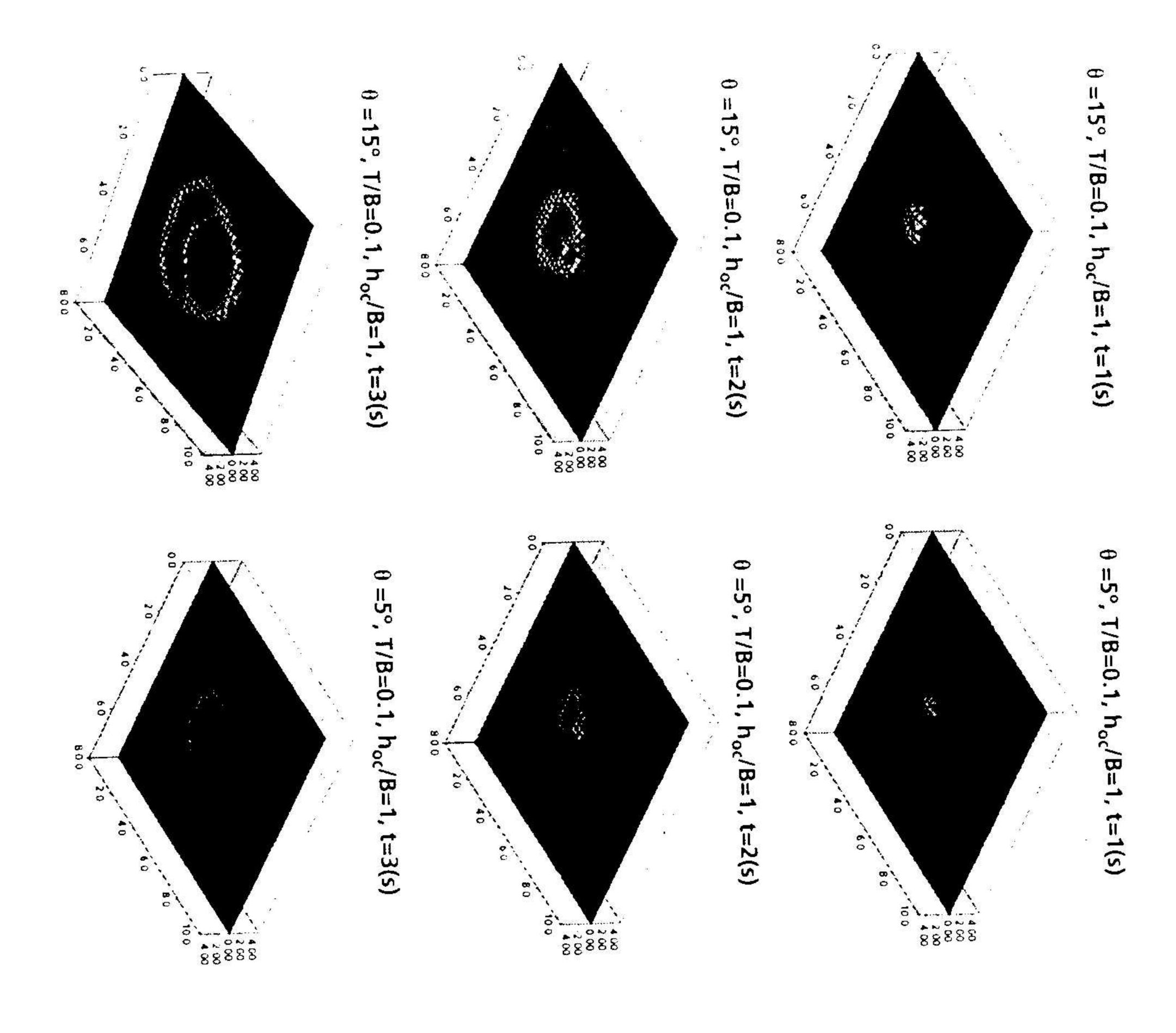


Figure 3: Numerical results for impulsive wave amplitude in depth integrated model to evaluate the influence of bed slope on the wave height and propagation pattern, in all cases the ratios T/B and h_{ac}/B is assumed 0.1 and 1; respectively, $C_d = 1$, $C_m = 1$, $\gamma = 1.85$, the wave heights multiplied for better display

4.1. Validation of prediction method

4.1.1 Comparison with experimental data

tion of experiments in each case is discussed briefly sented prediction method. dicted wave amplitude and the measured data are seen in the are used in the presented prediction method are determined in this Table. The descripsponding reference mentioned in Table 3. Moreover, landslide kinematics are listed. All of the experimental data are obtained from the correpresented in Table 3. In this Table, the sliding mass submerged sliding mass and the constraint of landsl the presented estimation method are selected from In this section, some available experimental measurements are fied. All of the experimental specifications and the procedure of prediction method are with laboratory data obtained from experiments in w sented prediction method. The estimated impulse wa As it is shown, the prediction error is about 5% and in measurements in several cases. The experim in Table 2. The comparison of preide density described later is satis hich the impulsive wave caused by ve amplitude is compared with labgeometry, Table ental data that are used to the basic dimensionless ratios that confirms a good reliability of pre 2. The comparison is made last columns of Table 3. failure mass density and used to validate the validate

of different shape of landslide in these cases, the present amplitude and comparison with experimental data is presented in Table 3. As seen, in spite and α_0 , thus the measured data of these kinematics parameters are used in prediction procesliding mass is a solid triangle block, equation (11) can noted that for laboratory works of Watts (1998) [31] and Heinrich (1992) [10], because the ry cases are presented in Table 2. The use of basic parameters metric parameters in these cases. The dimensions and other specifications of these laboratodure. Figure 4 shows also a schematic of these experiments and adefinition of landslide geouse of the prediction method in some real cases is discussed in the next sections. It must be used to apply presented prediction method for estimati real case condition of submerged landslide [30]. Thus tank is shown in Figure 4. The law of motion for underwater sliding block is similar to the ing of terminal velocity (u_l) and initial acceleration (α_0) and equation (11) can be used for the determination of In Grilli et al's experiments [6], the submerged sliding block is a semi-ellipse solid sheet on of impulse wave amplitude. The). A schematic slide kinematics parameters consistmethod works very not be used for determination of u_t in real cases, to predict the impulse wave equation (11) can be of their experimental well.

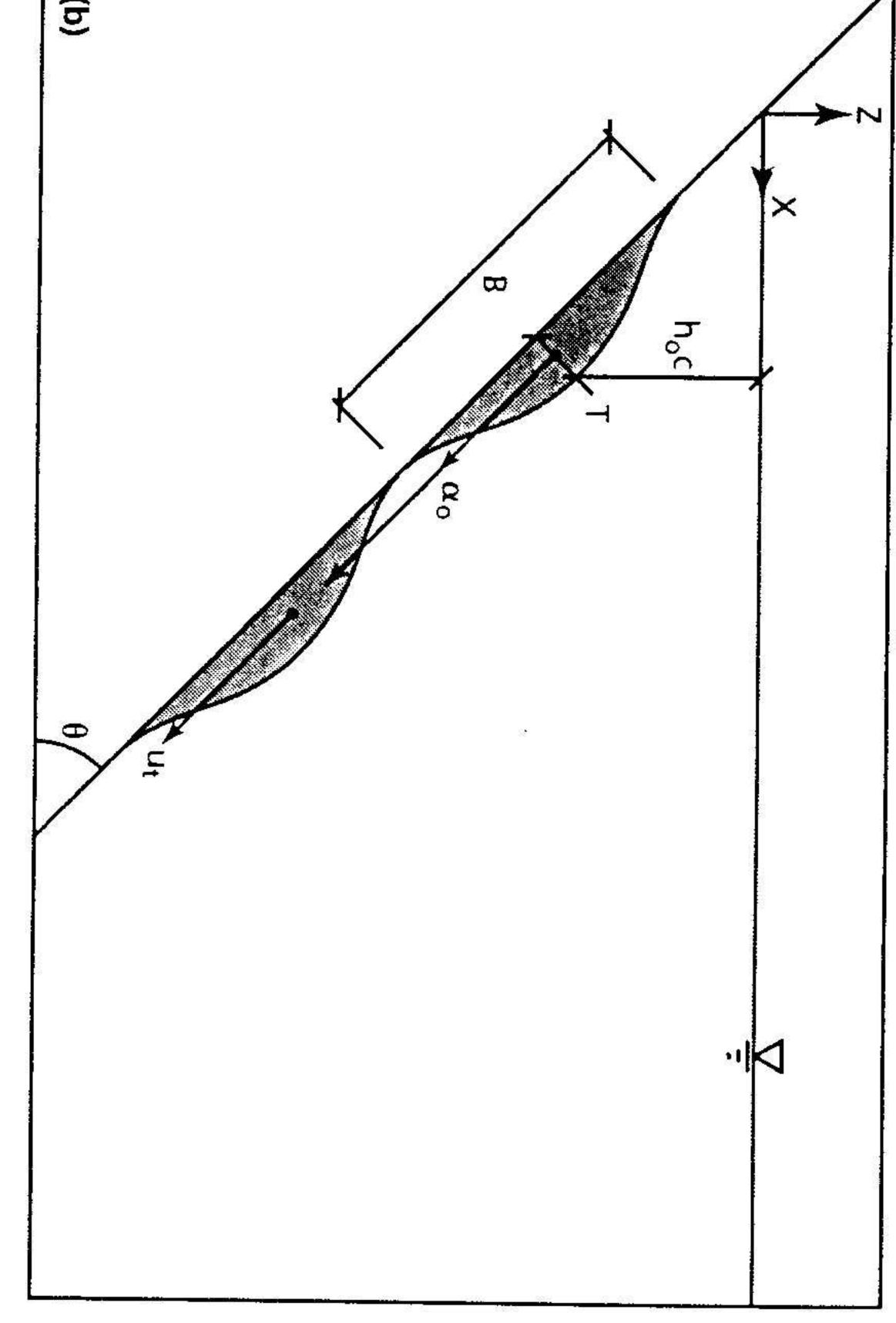
4.1. 2. Comparison with 3D numerical models

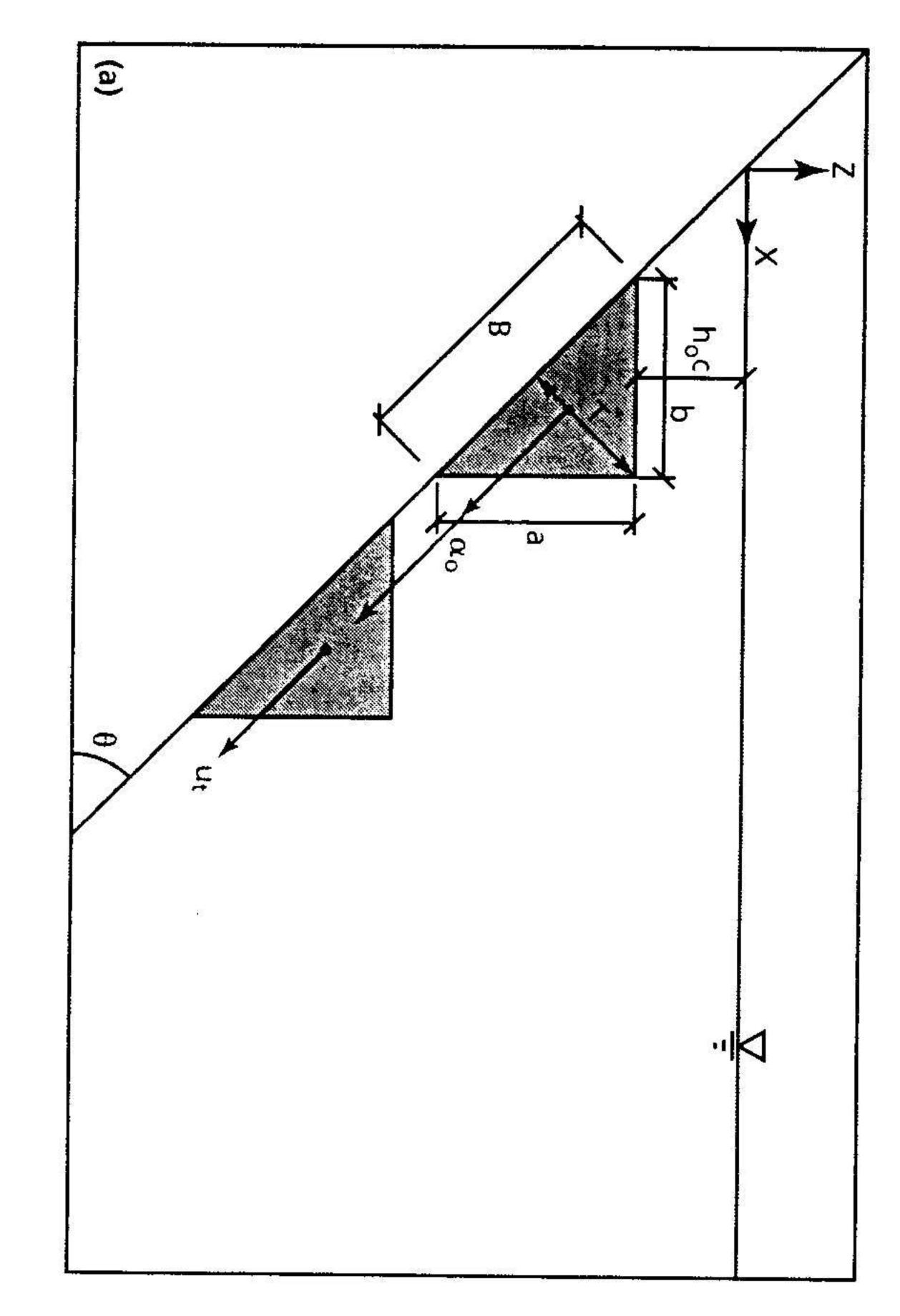
racy in simulating of a submerged landslide impulse as a three dimensional fully nonlinear potential flow model. All of the numerical cases properties are listed in In this section, presented prediction method is comp obtained from Watts et al. (2005) [33]. We use the r presented prediction method. The specification ared with umerical results wave is well documented [8, 9]. The BIEM The high degree of its accu-BIEM results are directly BIEM numerical mode are listed of BIEM to vali-3 Table 1.

Dam Engineering Vol XVII Issue

Ref	Reference	Landslide specifications					Water body Landslide spec. kinematr							Predicted Predicted ratio-fig 2 amplitude			6
no					C .	~	A CONTRACTOR) h _{Oc} (m)			- War (18 - 18 - 18 - 18 - 18 - 18 - 18 - 18	T/B	$h_{\theta\sigma} B$	a ₀ /\$0	$a_{0P}(m)$	a _{0M} (m)	90
		B(m)	T(m)		C _d					50°0	3.65	0.052	0.261	0.0014	0.0051	0.0052	1.9
6]	Grilli et al. (2002)		0.052	1.76	1.53	1.81	i mana	\$75009000AR00A0080)		1.95	0.21	0.5	0.611	0.02	0.0042	0.0045	6.6
31]	Watts (1998)	0.121	0.0608	0.81	1.67	2.1		**************************************	1		0.19	0.5	0.49	0.03	0.0057	0.0055	3.6
	Watts (1998)	0.121	0.0608	0.75	1.75	1.9	1		0.56	V7 (840)	CHECO CHECOLOGIC	######################################		20 10 <u>20</u> 0	0.22	0.22	0.1
[].] [10]		0.707	0.353	5 (4	%	2	45	0.01	0.6	1.5	0.24	0.5	0.014	0.72		AMERICAN VICTOR	<u> </u>

Table 3: Experimental validation of presented prediction method, the main specifications of experiments are mentioned in Table 2





No Landshde specifications Nater body spec Landshde kinematics Flack First Flack First							Water b	odv	Landslid	e kinema	atics	Basic ra	tios	Predicted	Predicted	2	Error
$ \begin{array}{ c c c c c c c c c c c c c c c c c c c$	No	Landstide specifications					v. a.c.							amplitude	amplitude		
$ \begin{array}{ c c c c c c c c c c c c c c c c c c c$							51/00							(fig 2)		(1) 1 (m)	0
1		D. Town IC C. V					B(dea)	$h_{\alpha c}(m)$	$u_i(m/s)$	$\alpha_0(m s^2)$	$S_0(m)$	$T \cdot B$	$h_{\theta C}/B$			*******	
1		B(m)		\subseteq_m				OK .			4.47	0.017	0.087	CRECRESE ES		1) 2002/00/20 (New Yorks)	* 194001 1
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1	2	1		1	,		- ' 	7000 100011	40 mm	0.255	4.47	0.017	0.17	0.0004	67 70 KN	26 XCSGC 20	80 (021)
1	3		***************************************	1	8 4 8		_) 	SI To any change exponential	83 AB B)		4.47	0.008	0.087	0.00034	#6000EC NO 0940E	90000400000000000000000000000000000000	All outs
5 1 0.013 1 1 1.83 3 0.007 1.067 0.255 4.47 0.025 0.087 0.0049 0.0049 0.005 2 7 1 0.052 1 1 1.85 15 0.5 1.84 0.757 4.47 0.052 0.625 0.0019 0.0085 0.009 5.5 8 1 0.052 1 1 1.85 15 0.625 1.84 0.757 4.47 0.052 0.625 0.00085 0.0039 0.0038 2.6 9 1 0.052 1 1 1.85 15 0.259 1.84 0.757 4.47 0.052 0.259 0.00062 0.0027 0.0029 4.5 10 1 0.025 1 1 1.85 15 0.259 1.84 0.757 4.47 0.025 0.0095 0.0042 0.0043 2.3 11 1 0.075 1 1 1.85	4	1	10 10 1001		1	50 000 at	, <u>,</u>	9 555 8 70 555	60 ACCASON NO. 150.		4.47	0.013	0.087	0.00053	0.0024	3	4
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	1	8 1	0.2		1	1.85	30	0.5	2.55	1.40	7.7/	0.2					<u>l</u>

Table 4: Comparing presented prediction method with BIEM numerical model results [8] as a fully nonlinear potential flow model, the main specifications of BIEM numerical model are mentioned in Table 1

comparison of results can be seen in the last columns in dimensionless ratios that is used in the prediction method are listed in Table 4. slide has a maximum deviation from 3D numerical model as 5% and it confirm an numerical results. The predicted impulse wave amplitude shown, the results of simple prediction method are in wave amplitude in a water body. excellent reliability of presented engineering estimation method to predict the impulse Table 4 shows the comparison of BIEM and presented The details of the prediction procedure and the basic this table. caused by under water landgood agreement with BIEM prediction method.

4.2. Application of prediction method in some real cases

evaluation of prediction method it is applied in some real cases of submarine LGWs. checked in different condition by experimental and other numerical data, presented estimation method is developed based on dam reservoir related to the tsunamis which were caused by submarine Most observations and real case studies about underwater landslide generated waves mass failure. Thus, although the characteristics and for additional

4.2.1. The 1946 Unimak, Alaska submarine landslide

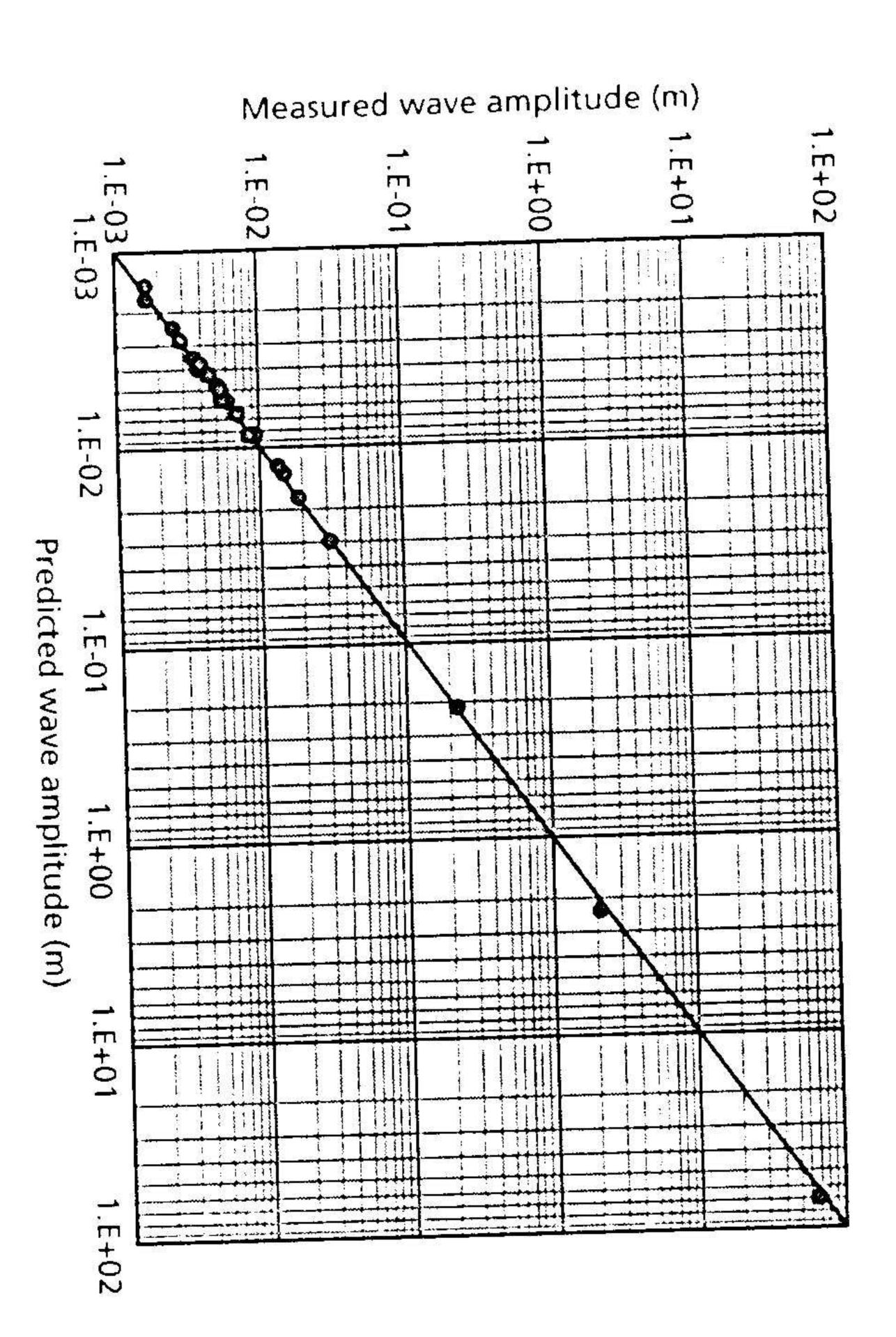
although the bed slope in the real case is $\theta=4.3^{\circ}$ be mentioned that the estimation graph used in this case The comparison shows acceptable reliability of prediction using equation (11) for determination of S_0 , the impulsive wave amplitude estimated T/B=0.0075 and $h_{0C}/B=0.042$ is $a_0/S_0=-0.0004$. Assuming that the shape related coefficients of sliding mass consist of added mass and drag coefficients are both equal to one and ness is 300m. The initial submergence depth is 1600m. The results come from Figure 2 for 71.6m can be compared with other numerical models which result in the value of 64m [32]. submarine failure mass can be defined by the parameters presented here [32]: the bed slope shaking of the initial earthquake. Based on the results of investigations, the geometry of believe that the waves were caused by a submarine lands and kept going to Antarctica. The April 1, 1946 wave that killed 159 people in Hawaii, smashed into the More than half a century later, scientists are beginning to piece together the puzzle , the sliding soil density is 1.85, the length along the bed is 40,000m and the thickoutcome of investigations is that the most scientists lide, presumably, triggered by the method presented here. It must is according to bed slope Marquesas Islands

4.2.2. The 1994 Skagway, Alaska submarine landslide

with the geometric and kinematics of them listed here [32]: For slide A; the angle of bed at the head of Taiya Inlet. The impulsive waves could be caused by two distinct landslides water landslide formed during the collapse of a cruise The impulse wave of November 3, 1994 in Skagway, Alaska, was generated by an under-The initial submergence is 24m and the shape coefficients , the soil density is 1.85, the length and thickne ship wharf undergoing construction ss are 180 and 20m, respectively. are unit. It can be obtained that

edicted and other numerical result characteristic 0.023obtained numerical dimensionless slide the obtained that can method can bed slope ₿; bc length submergence bed angle acquired ratios; and the S amplitude considered that the is less than the sliding block ō obtained and 0.44, an estimated value is wave can be 0 acceptable amplitude S respectively. bed slope in F determined 215m and its thickness is 5m. bed reliability = slope acceptable Using omparing present-

less than prediction graph



Conclusion

tial submergence of sliding block diminishes. slides. These effects are fully nonlinear particularly increasing effect of sliding thickness on the sliding mass. Furthermore, the influence of bed slope mass is assumed S_0 , defined based on the initial acceleration and terminal velocity of wave amplitude, wave length and its propagation slope (B) (i.e. T/B), the initial submergence of center as follow: the maximum thickness of sliding mass ing mass length (B) (i.e. h_{0C}/B). The main characteristic related to kinematics of sliding cluded that the main geometrical ratios which affect acteristics of submerged failure mass are investigate voir using a higher order Boussinesq-type numerica A sensitivity analysis carried out for submerged landslide wave height multiplied in shallower landfor steeper bed slope when the inipattern. d in dimensionless form. It is conof failure (T) over its length along the bed is investigated on the generated wave impulse The mass (h_{0C}) over results wave amplitude are geometric show in dam impulse

engineering utility for the first estimate of LGWs. and a good agreement is obtained. impulse wave amplitude is appraised and compared with the predictive method is applied in some real cases ed values is about 5% and the reliability of prediction method is confirmed. In addition, presented. The method is validated using several experimental data and well-known 3D numerical models, and good agreements were obtain-Lastly, a simple and applied approach for estimate The present method can be used as an applied and submarine mass failure The maximum error of of impulse wave well-known numerical models amplitude and the

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University of Technology, Iran. Power of the Islamic Republic of Iran is appreciated Financial assistance was also provided partly by The financial support of Water Resources Manag Organization, number: Office DAM1-84010). Ministry of Sharif

- = the water surface displacement from still water level scaled by a0
- = the slide mass density [MT²L²]
- θ S = the bed slope angle [deg.] the nonlinearity ratio = a_0/h_0 [-]
- Ŧ d = the horizontal gradient vector = $(\partial/\partial x, \partial/\partial y)$ = the frequency dispersion ratio = h_0/l_0 [-]
- α_0 = the initial acceleration of slide mass [LT⁻²]
- the weighting parameter for determination characteristic depth
- a two-component vector which is used in derivation (Equation

æ (°° 5 € h_{0C} h(x,y)the impulse wave amplitude [L] the length of slide mass along the bed slope [L] the characteristic water depth [L] Drag coefficient [-] Added mass coefficient [-] the depth of moving bottom boundary from still water level scaled by h_0 [-] the horizontal wavelength scale [L] Initial still water depth at center point of sliding mass [L] Acceleration due to gravity [LT-2]

 \mathcal{B}

- 0 $\overline{}$ the water pressure scaled by $\gamma_w a_0$ [-]
- .S. the kinematics length sacale of sliding mass (Equation 11)[L]
- dimensionless time scaled by $l_0/(gh_0)^{1/2}$ [-]
- the maximum thickness of the slide mass [L]
- the vector of horizontal velocity components (u, v) scaled by $\varepsilon.(gh_0)^{1/2}$ [-]
- 100 terminal velocity of slide mass [LT]
- u_0, u_1, u_2 the dimensionless factors used for expanded form of u (Equation 7) [-] the velocity component in vertical direction scaled by $(\varepsilon/\mu).(gh_0)^{1/2}$ [-] the dimensionless factors used for expanded form of w (Equation 8) [-]
- . the horizontal coordinates scaled by I0 [-]
- z_a, z_b them, scaled by h0 [-] the elevations that the horizontal velocity components are described in
- 1 water depths, the characteristic variable z_a, z_b [-] depth which is a weighted average of two distinct
- 1.1 the vertical coordinate scaled by h_0 [-]

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